

A review of simulated climate change impacts on groundwater resources in Eastern Canada.

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Manuscripts

1 **A review of simulated climate change impacts on groundwater**
2 **resources in Eastern Canada**

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13

14 **Abstract**

15 In Eastern Canada, groundwater is the main water supply for most of the rural regions
16 and in many large urban communities. An understanding of the impacts of climate
17 change on this resource is crucial for sustainable water management in this region. The
18 objectives of this paper are to summarize the state of knowledge about possible climate
19 change impacts on groundwater dynamics in Eastern Canada thus providing a clearer
20 understanding of future conditions. Twenty-two studies are reviewed including Ontario,
21 Québec, New Brunswick, Nova Scotia, and Prince Edward Island, to identify the impacts
22 on groundwater recharge and river baseflows. The studies disagree in their estimates of
23 changes in future recharge conditions, and no trend from West to East was revealed. This
24 could be due to the use of different modelling approaches (model type, climate change
25 scenario, future horizon). Nonetheless, more inter-annual variability during the summer
26 and earlier snowmelt periods causing seasonal shifts in the recharge cycle are expected.
27 This review provided new insights that lead to the following recommendations for future
28 studies: 1) use a variety of climate models and emission scenarios; 2) promote the use of
29 integrated models when possible; 3) study long-term climate change impacts on
30 groundwater resources at different scales; 4) simulate the combined effects of climate
31 change and other pressures; and 5) develop models that cover other regions of Eastern
32 Canada as dictated by stakeholders and water managers.

33

34 **Key words:** Groundwater, Eastern Canada, Model, Climate change

35

36 **Résumé**

37 Dans l'est du Canada, l'eau souterraine est la principale source d'approvisionnement en
38 eau en région rurale et dans plusieurs grandes villes. La compréhension des impacts des
39 changements climatiques sur cette ressource est cruciale pour la gestion durable de l'eau
40 dans cette région. Les objectifs de cette revue étaient d'établir l'état des connaissances
41 sur les impacts possibles des changements climatiques sur la dynamique des eaux
42 souterraines dans l'est du Canada afin de mieux appréhender les conditions futures.
43 Vingt-deux études portant sur les provinces de l'Ontario, du Québec, du Nouveau-
44 Brunswick, de la Nouvelle-Écosse et de l'Île-du-Prince-Édouard ont été analysées pour
45 déterminer les impacts sur la recharge des eaux souterraines et les débits de base des
46 cours d'eau. Les études ne s'entendent pas sur une l'évolution future de la recharge et
47 aucune tendance de l'ouest vers l'est n'a été mise en évidence. Ceci pourrait être causé
48 par les différentes approches de modélisation utilisées (type de modèle, scénario
49 climatique, horizon futur). Néanmoins, les résultats montrent généralement une plus
50 grande variabilité interannuelle de la recharge estivale et une recharge printanière plus
51 hâtive. Les études montrent que les débits d'étiage pourraient diminuer, ce qui induirait
52 des conditions critiques pour les approvisionnements en eau et pour les écosystèmes.
53 Cette revue a fourni un nouvel éclairage qui a conduit aux recommandations suivantes
54 pour les études futures: 1) utiliser une variété de modèles climatiques et de scénarios
55 d'émission; 2) utiliser des modèles intégrés lorsque possible; 3) étudier les impacts à long
56 terme des changements climatiques sur les ressources en eaux souterraines à différentes
57 échelles; 4) simuler les effets combinés du changement climatique et d'autres pressions;

58 et 5) élaborer des modèles qui couvrent d'autres régions de l'est du Canada dictées par les
59 intervenants et les gestionnaires de l'eau.

60

61 **Most-clés:** eau souterraine, est du Canada, modèle, changement climatique

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64 INTRODUCTION

65 Throughout the world, climate change has become evident from the observations of
66 global mean temperature rise and the higher frequency of extreme weather events (IPCC
67 2013). These changes could affect groundwater resources in a variety of ways (Taylor et
68 al. 2012). In a warmer climate, increasing demand of water is expected to lead to
69 increased pumping rates for drinking water and irrigation. Conditions of depleting
70 groundwater reservoirs and reduced groundwater flow to rivers, wetlands, and lakes are
71 either already clearly evident or expected around the world in the next decades (Ferguson
72 and Gleeson 2012; Crosbie et al. 2013; Castle et al. 2014; Döll et al. 2014; Kløve et al.
73 2014; de Graaf et al. 2017).

74 Canada is amongst the world's richest countries in terms of fresh water resources
75 (FAO 2003). Throughout Canada, groundwater provides approximately 30% of the
76 population with potable water and is often the only source of water in rural areas (ECCC
77 2017). When investigating the presence of climate change impacts in past time series of
78 baseflows and groundwater levels across the country, Rivard et al. (2009) have identified
79 mixed increasing and decreasing trends country-wide. However, the Atlantic Provinces
80 showed statistically significant decreases of baseflows, while some regions north of 55°N
81 had increasing baseflows. The majority of the longer time series showed negative trends
82 during the summer months. Chen (2015) assessed the impact of climate change on
83 groundwater dynamics at the continental-scale using the integrated HydroGeoSphere
84 model (Therrien et al. 2010) for the entire country and the northern part of the United-
85 States. Confirming some of the trends identified by Rivard et al. (2009), Chen (2015)

86 showed that annual streamflow in most of the northern main rivers could increase under
87 climate change, but results for the southern regions were less clear.

88 In 2016, more than 68% of the country's population resided in Eastern Canada,
89 notably in Ontario and Québec (Statistics Canada 2016). In Eastern Canada, groundwater
90 is a frequent source of water for many usages like municipal water supply, irrigation for
91 agriculture, and private wells for potable water. In Ontario, Québec, and the Atlantic
92 Provinces respectively, 14, 17, and 12% of irrigation water for farming was taken from
93 groundwater in 2012 (Statistics Canada 2012). A rainfall rate increase is expected, based
94 on a high-resolution regional climate modelling study over the Great Lakes basin
95 (d'Orgeville et al. 2014). Among other effects, climate change is expected to lead to
96 earlier spring flood events and to more severe summer low flows (Ouranos 2015). In the
97 Atlantic Provinces, more storm events and increasing storm intensity, rising sea levels,
98 storm surges, and coastal erosion are expected from climate change (Climate Action
99 Network 2017). The impacts of climate change on groundwater resources are not fully
100 understood and a global picture of possible future conditions is not yet available for water
101 managers.

102 The simulation of possible groundwater flow conditions provides interesting outlooks
103 for potential conditions in future decades. However, modeling studies are expensive, are
104 often of local application, and can be based on a wide array of future conditions (e.g.,
105 stemming from different climate models, groundwater flow models, and emissions
106 scenarios). Holman et al. (2011) made methodological recommendations aimed at
107 improving the assessment of climate change impacts on groundwater. In recent years, the
108 general approach has tended towards implementing these recommendations. As a result,

109 recent studies now commonly used multiple climate change scenarios, including
110 scenarios from global climate models (GCMs), regional climate models (RCMs),
111 different greenhouse gas (GHG) emission scenarios, different downscaling methods and
112 different future horizons (see discussion in Kurylyk and MacQuarrie 2013). Additionally,
113 modeling studies are performed on a wide variety of spatial scales, using models that
114 simulate the dynamics of the entire water cycle or fluxes in some reservoirs only (soil,
115 aquifer, or river). For all these reasons, it is difficult to compare modeling results and to
116 use them in water management decisions that would include adaptation options. The
117 objective of this paper was to summarize the state of knowledge about simulated climate
118 change impacts on groundwater dynamics in Eastern Canada, in order to provide a clearer
119 understanding of possible future conditions and to make recommendations for future
120 studies.

121

122 **BACKGROUND INFORMATION**

123 *Hydrological, hydrogeological, and integrated models*

124 A wide array of models are available to study issues related to groundwater flow.
125 Models such as the spatially distributed HYDROTEL model (Fortin et al. 2001) or the
126 conceptual HBV model (Lindström et al. 2007), can be used to simulate surface flow
127 processes. They usually include a highly simplified representation of the aquifer that
128 empties with a calibrated reservoir coefficient. Nevertheless, surface flow models can be
129 useful to study climate change impacts on groundwater resources if they are used to
130 provide information on future base flow which corresponds to groundwater discharge to
131 the surface in rivers.

132 Groundwater flow models such as MODFLOW (saturated flow; Harbaugh 2005) and
133 FEFLOW (saturated and unsaturated flow; DHI-WASY 2013) are widely used. This
134 category also includes models dedicated to simulate recharge and which are used with
135 groundwater flow models to provide the upper boundary condition of water that reaches
136 the aquifer through the surface. The most widely used in this category is the HELP model
137 (Schroeder et al. 1994). The SUTRA model (Voss and Provost 2002) allows the
138 simulation of coupled groundwater flow and groundwater temperature and is used to
139 understand the warming effect of climate change on groundwater flow systems. The
140 SEAWAT model (Guo and Langevin 2002) is software based on MODFLOW/MT3DMS
141 (Bedekar et al. 2016). It simulates 3D density-dependent groundwater flow and is
142 frequently used to simulate saltwater intrusion in coastal aquifers.

143 Integrated models represent the entire land water cycle and usually include
144 evapotranspiration, snow accumulation and melting, runoff, water routing at the surface
145 and in the river channel, infiltration, groundwater flow and groundwater discharge to
146 surface reservoirs. HydroGeoSphere (Therrien et al. 2010), MikeSHE (DHI Software
147 2007), and CATHY (Camporese et al. 2010) are three examples of integrated models that
148 are widely reported in the scientific literature.

149

150 ***Emission scenarios, climate models, and projected horizons***

151 Earlier climate modeling studies were based on GHG emission scenarios. IPCC (2000)
152 recommended the use of different emission scenarios to be used in climate modeling
153 studies (A1, A2, B1 and B2, including sub-scenarios). Representative Concentration
154 Pathways (RCPs) including four GHG concentration trajectories are now used to replace

155 the former emission scenarios in climate modeling and research (van Vuuren et al. 2011).
156 The RCPs describe potential futures of the main drivers of climate change: greenhouse
157 gas and air pollutant emissions. The scenarios are named based on the change in radiative
158 forcing in 2100 compared to pre-industrial values. The RCP4.5 scenario (optimistic)
159 represents an increase in radiative forcing of 4.5 W/m^2 relative to pre-industrial values. It
160 is associated with a capping of emissions which would stabilize the radiative forcing
161 caused by climate change in 2100. The RCP8.5 scenario (pessimistic) represents no
162 change in current human behaviour. Emissions continue to rise beyond 2100 when the
163 radiative forcing is increased by 8.5 W/m^2 relative to pre-industrial values. Most of the
164 recent climate change impact studies use these two RCPs to account for GHG
165 concentration scenarios.

166 Climate projections are generated either directly from GCMs or from RCMs driven by
167 GCMs. The shift from SRES to RCPs also corresponds to the replacement of the CMIP3
168 global climate model ensemble (Meehl et al. 2007) with the CMIP5 ensemble (Taylor et
169 al. 2012). There are currently 20 GCMs and 12 pairs of RCM-GCM combinations for
170 North America operated by different research groups around the world for which climate
171 projections are available. Using multiple climate scenarios combining climate models and
172 emission scenarios is the usual approach for climate change impact studies to provide an
173 array of possible futures.

174 Data from climate models that are readily available for use in hydrological and
175 hydrogeological models are air temperature (daily minimum, maximum or average) and
176 precipitation. These data are generated on grids of various sizes, depending on the
177 climate model used to simulate them. Because the hydrological and hydrogeological

178 models are usually of more local scale than the climate models' resolution, and their
179 outputs usually include statistical bias, the climate model outputs are often adjusted using
180 post-processing methods. When using statistical downscaling (Thiemeßl et al. 2010;
181 Teutschbein and Seibert, 2012), climatic variables are linked to local meteorological
182 variables using different methods (e.g., quantile mapping, linear regression, analog
183 method). This method takes into consideration changes in climate variability. The delta
184 change method (Diaz-Neto and Wilby 2005) calculates mean deviations in temperature
185 and precipitation between future and past periods. It can be used on monthly or daily
186 data. This method assumes that rainfall frequency is not modified with climate change.
187 Dynamical downscaling (Giorgi and Mearns 1991) corresponds to running a fine scale
188 regional climate model on a sub-domain driven at its boundary by a global climate
189 model. Results from these models sometimes still need to be post-processed.

190 The earliest studies of climate change impacts on groundwater resources used climate
191 scenarios for 30 to 40 years over which the future conditions were to be considered
192 constant, and compared to simulated recent climate. The more recent climate impact
193 studies use continuous scenarios in which the simulated future conditions evolve through
194 time from past conditions in the 1950s to future 2100 conditions.

195

196 **STUDIED REGIONS**

197 The surveyed literature concerns studies that simulate future groundwater flow
198 conditions for the provinces of Ontario, Québec, New Brunswick, Nova Scotia, and
199 Prince Edward Island. Newfoundland and Labrador was not included in this overview
200 because no provincial- nor local-scale studies were available. Only peer-reviewed studies
201 in journal papers and theses reporting the use of groundwater flow models, integrated

202 surface-groundwater flow models, recharge models, as well as surface flow models
203 which include at least a simple representation of aquifers and a quantification of
204 baseflows were considered. Surface flow studies with little or no representation of
205 groundwater reservoirs or without any explicit quantification of recharge or groundwater
206 contributions to rivers were excluded. A total of 22 studies have thus been summarized
207 spanning a wide region between southern Ontario and the Magdalen Islands (QC)
208 (Figure 1).

209

210 ***Ontario-based studies***

211 Six studies focusing on groundwater resources and climate change have been reported
212 in the literature for Ontario (see Table 1 and Figure 1b). The studies have been conducted
213 only in southwestern Ontario with no models applied to eastern or northern Ontario.
214 Various hydrological modelling approaches, from strictly recharge models (HELP), to
215 fully-integrated models (e.g., HydroGeoSphere (HGS) and MikeSHE) have been
216 employed. Most of the studies were based on data from multiple climate models. The
217 emission scenarios included A2, B2, and a doubled CO₂ equilibrium (prior to the IPCC
218 (2000) Special Report Emission Scenarios). Projected simulated durations ranged from
219 20 years (Sultana and Coulibaly 2011) to 60 years (Brouwers 2007). The earliest future
220 horizon starts in 2020 (Brouwers 2007), and the latest ends in 2080. Some studies
221 projected a future time span without using specific years (40 years in Jyrkama and Sykes
222 2007 and Colautti 2010). McLaren and Sudicky (1993) projected their future steady-state
223 flow conditions in 2050 (no time-span).

224 One of the most studied aquifer systems in Canada regarding climate change impacts
225 on groundwater at the regional-scale is the Grand River watershed (6,800 km²) in

226 southern Ontario. As early as the 1990s, researchers were anticipating the potential
227 impacts of the climate change on its groundwater resources. For instance, McLaren and
228 Sudicky (1993) used a 2D, steady-state groundwater flow model for a subregion of the
229 Grand River watershed to examine predicted head and baseflow changes. In the
230 groundwater model, a recharge change of -15 to -35% lead to baseflow changes of -17 to
231 -39%.

232 Jyrkama and Sykes (2007) presented a physically-based method to evaluate temporal
233 and spatial variability of climate change impacts on the recharge over the Grand River
234 watershed (6,800 km²). They used a GIS version of the HELP model, a distributed water
235 balance and routing model using a pseudo-2D representation to simulate the recharge.
236 Combining scenarios for temperature, precipitation and solar radiation changes yielded
237 +10% to +53% for future groundwater recharge, -12% to +10% for future surface runoff,
238 +3% to +12% for future evapotranspiration. Among various conclusions for the Grand
239 River watershed, the authors outlined that the increased recharge projected for the future
240 would not be uniformly distributed. Moreover, impacts were controlled by local
241 groundwater elevations, types of soil and land uses.

242 Brouwers (2007) coupled the HELP model with the HydroSphere model (precursor of
243 the HGS model) for saturated groundwater flow to simulate the projected behaviour of
244 the Alder Creek basin (80 km²), a subwatershed of the Grand River. The simulations
245 showed a shift in the snowmelt timing, causing a general reduction of the runoff, an
246 increase of the evapotranspiration (mainly during the summer months), and a general
247 increase of the infiltration. Impacts on groundwater were generally lower than on surface
248 water. Changes in average monthly recharge ranged from +0.36 mm (urban land, 2040-

249 2060) to +4.12 mm (agricultural land, 2060-2080). The author reports that recharge
250 increased the most during the spring, with a shift of the melting season towards earlier
251 dates, although this is not apparent from the average monthly values.

252 Colautti (2010) applied the integrated HGS model in the Grand River watershed
253 (6,800 km²) using five climate scenarios. Future scenarios were constructed based on
254 modifying the 1960-1999 historical precipitation record (by -5%, +5%, +10%, +15%,
255 +20%), bounded by GCM-based climate scenario ranges. These scenarios yielded
256 recharge changes of -5% to +22%, and river discharges changes of -15% to +59% with
257 changes in groundwater levels between -0.55 m and +1.25 m. At the local-scale, Colautti
258 (2010) highlighted that simulated depth below ground surface to water table responded
259 differently from one zone to another, suggesting that the local flow patterns may be more
260 sensitive to future climate changes.

261 Sultana and Coulibaly (2011) used the integrated MikeSHE model for the evaluation
262 of climate change impacts on the Spencer Creek watershed (291 km²) in southern Ontario
263 The simulations predicted a change of annual groundwater recharge between -6% and -
264 0.5%. This decrease was particularly visible during summer and fall seasons. The authors
265 did not quantify the change in river baseflow, but they note that it is expected to decrease
266 due to the reduction of recharge.

267 Recently, Motiee and McBean (2017) applied HELP in the Guelph region (unspecified
268 study area) of the Grand River watershed to investigate the impacts of climate change on
269 recharge. The authors predicted increased evaporation and decreased summer recharge,
270 with increased recharge in the winter months due to changing freeze/thaw dynamics.
271 They determined a future recharge change between +7% and +12%.

272

273 ***Quebec-based studies***

274 Ten local to regional-scale climate change and groundwater studies in the province of
275 Québec have been reported in the literature (see Table 2 and Figure 1c). The models used
276 a surface flow (HYDROTEL), a groundwater flow (MODFLOW), or integrated surface-
277 groundwater flow (CATHY, HGS, and MikeSHE). All the studies were based on data
278 from multiple climate models. Simulated future time periods are all similar in lengths,
279 ranging between 24 years (Sulis et al. 2012) and 32 years (Sulis et al. 2011), with
280 projected horizons starting as early as 2010 (2010-2039; Quilbé et al. 2008) and as late as
281 2071 (2071-2100; Cochand 2014). The latest end for future horizons was 2100 (Cochand
282 2014). The Québec-based study from Lemieux et al. (2015) in the Magdalen Islands is
283 discussed in the Atlantic studies because of their geographical proximity.

284 Quilbé et al. (2008) assessed the effect of climate change on the Chaudière River
285 watershed (6,682 km²), with the HYDROTEL model. The results that are of interest for
286 the current paper concern the critical streamflow sequences over seven days and return
287 periods of two and ten years (Q_{2-7} and Q_{10-7}), as well as critical streamflow sequences
288 over 30 days for five years return periods (Q_{5-30}), which are considered to represent river
289 baseflows during the summer period. Interestingly, the scenarios performed with the
290 Delta method showed changes in the critical low flows of -23% to -5% for the Q_{2-7} , -25%
291 to -7% for the Q_{10-7} , and -29% to -7% for the Q_{5-30} . The simulations performed with
292 statistical downscaling showed no obvious effect of climate change on summer low
293 flows.

294 Sulis et al. (2011) evaluated the climate change impacts on the Anglais River
295 (690 km²; Montérégie region) with the integrated CATHY model. The authors identified
296 that the impacts on the river flow at the outlet were greater during winter peaks and
297 summer droughts. In future conditions, the total recharge changed by +16%; however, the
298 changes throughout the year were not constant. The winter recharge was higher (+49%)
299 due to increased rain and snowmelt, the spring recharge remained the same, and the
300 summer recharge was lower (-8%) due to increased evapotranspiration. The fall recharge
301 was higher (+22%) due to increased precipitation. A spatial analysis of recharge patterns
302 shows that the greatest variations in recharge are expected to occur at the highest
303 elevations. The future simulated river flows at the outlet during the June, July, and
304 August summer months (considered as baseflows) were 56% lower than those of the
305 reference period.

306 In the same des Anglais watershed, Sulis et al. (2012) also investigated the impact of
307 climate change uncertainty in hydrological processes using the CATHY model.
308 Hydrological responses (streamflow, recharge and groundwater storage) reacted
309 differently to the precipitation and temperature variations between the climate models.
310 River discharge changes varied from -18% to +11%. For the future climate, the low flow
311 occurrence frequency increased for all simulations, with a percentage change in low flow
312 occurrence between +12% and 25% for the members (comparison based on the
313 percentage of days with a discharge lower than the first decile of present-day discharges).
314 The change in total recharge varied between a -15% and +4%. The authors underline the
315 impact of changing sequences of rainy days on groundwater recharge.

316 Bourgault et al. (2014) simulated aquifer-peatland-river interactions under climate
317 change in the Lanoraie peatland complex (364 km²). The authors used MikeSHE and
318 activated only its groundwater flow component. From three climate scenarios, the
319 recharge was estimated using a simple water budget calculation, based on the hypothesis
320 that there is no runoff on the highly permeable sand aquifer. The resulting recharge
321 variations ranged between 0% and -50%. The authors used average recharge scenarios of
322 -50% and -20% to simulate groundwater flow. It appeared that the storage capacity of the
323 organic deposits contributed to prevent drastic drawdown to the surrounding aquifers and
324 limit river baseflow decreases. However, despite the mitigating role of the peatland, the
325 climate scenarios indicated a change in river baseflows between -41% and -16%.

326 Cochand (2014) simulated the impact of climate change on the Saint-Charles River
327 watershed (553 km²) using the integrated HGS model. The summer and fall low flows
328 (June to October) changed between -25% and -10%, mainly due to increased
329 evapotranspiration. Increasing temperatures in winter and early snowmelts triggered
330 recharge increases (+150 to +300%) between December and March and recharge
331 decreased (-50%) in April-May. Recharge decreased also between June and October, but
332 less markedly (-4 to -8%). The winter recharge increase was observed in the higher
333 topography areas.

334 Levison et al. (2014a) developed a groundwater flow model for the Covey Hill Natural
335 Laboratory in southern Québec (173 km²; Montérégie region) using a steady-state
336 groundwater flow model (MODFLOW). Similarly to Bourgault et al. (2014), the authors
337 compared the net precipitation from the future horizon to that of the reference period for
338 all the members, with changes in net precipitation ranging between -30% and +10%.

339 Assuming that recharge variations should be similar to the net precipitation variations,
340 they imposed recharge change scenarios of -30%, -15% and +10%,. The total
341 contribution from the aquifer to the rivers and streams over the study area varied between
342 -44% and +14% for the +10% and -30% recharge scenarios respectively. Near the
343 peatland, the direction of hydraulic gradients also changed in the future conditions,
344 making the peatland feed the aquifer during the summer, the fall and the winter seasons.

345 In the same location, Levison et al. (2014b) simulated the dynamics of small bedrock
346 springs under climate change conditions using the HGS model representing local scale
347 discrete fractures. Levison et al. (2014b) applied 10 climate change scenarios directly to
348 the HGS model instead of using a sensitivity analysis approach and imposed net
349 precipitation at the soil surface. The simulated spring flow rates changed by +5 to +6%
350 under future conditions, depending on the spring elevation. Results also indicated a
351 significant increase in the number of days of spring flow activity (+1 to +2%) and
352 generally more variability in the duration of the flow for all springs, although these
353 increases varied considerably depending on the spring location. The authors proposed
354 that this location might be resilient enough to face the projected climate changes, but
355 emphasized the importance of improving follow-up programs on ecologically sensitive
356 sites to acquire more information about the ability of natural habitat to face different
357 climate conditions.

358 With the Hydroclimatic Atlas, the CEHQ (2015) synthesized the impact of climate
359 change on the river regime of 50 medium size basins (500 to 20 000 km²) in southern
360 Québec. The HYDROTEL model was used to simulate flows. Changes in a series of
361 indicators were quantified to assess the impact of climate change on the watershed

362 hydrology. The indicators of interest here are those for summer low flows which are
363 associated with baseflows. The results showed that summer low flows events will
364 generally be more severe and longer for all scenarios and all watersheds: Q_{2-7} changed by
365 -54% to -6%, Q_{10-7} changed by -63% to -9%, and Q_{5-30} changed by -56% to -8%.

366 Levison et al. (2016) also investigated long-term trends in groundwater recharge and
367 discharge for the Covey Hill Natural Laboratory using the MODFLOW model. The
368 authors compared the simulated responses of the groundwater system using observed and
369 simulated data from 1900 to 2010 and also used projected climate data for 2041-2070.
370 They used the Levison et al. (2014a) calibrated model, but applied the 10 RCM climate
371 change scenarios from Levison et al. (2014b). It was assumed that recharge variations
372 should be similar to the net precipitation variations (-4% to +15%). For the future
373 scenarios, average baseflows for the three rivers increased (from +10% to +14%), and
374 average spring flow also increased (+18%). These results, in comparison to Levison et al.
375 (2014a; 2014b) illustrate the complexity and uncertainty for making predictions for
376 groundwater and climate change.

377

378 *Atlantic Canada studies*

379 Thus far in the literature, six climate change related groundwater modelling studies
380 have been reported for New Brunswick, Nova Scotia, and Prince Edward Island.
381 Although it is located in the Province of Québec, the Lemieux et al. (2015) study in the
382 Magdalen Islands is reported here because of its geographical proximity. The models
383 simulate recharge (HELP), groundwater flow and water temperature (SUTRA),
384 unsaturated-saturated groundwater flow (FEFLOW), and salt water intrusion

385 (SEAWAT). The emission scenarios are A1B, A2, B1, and B2. The future model runs
386 span 27 years (Lemieux et al. 2015) to 90 years (Green and MacQuarrie 2014). The
387 future horizons start as early as 2011 and end as late as 2100 (both from Green and
388 MacQuarrie 2014) (see Table 3 and Figure 1d).

389 Hansen (2012) used the SEAWAT model to estimate submarine groundwater
390 discharge under climate change conditions on the coastal aquifer of Summerside (PEI).
391 The authors simulated a combination of changes in recharge (assumed to be the same as
392 changes in precipitation), changes in sea water levels and changes in aquifer pumping
393 rates. Their results show that climate change has limited impact coastal groundwater
394 discharge in the study area. Sea-level rise leads to significant saltwater intrusion (between
395 30 and 60 m) but is mitigated by a +5 to +8% change in recharge. Increasing groundwater
396 pumping appears to be the dominant process for saltwater intrusion.

397 Kurylyk and MacQuarrie (2013) simulated the impacts of climate scenarios on the
398 recharge with the HELP model at the local-scale watershed, on the Otter Brook
399 watershed (NB) (9.5 km²). The annual recharge changes varied between -6% and +58%.
400 The authors show that the post-processing method had a large impact on the results,
401 sometimes a larger impact than the emission scenario.

402 Kurylyk et al. (2014) applied the seven recharge scenarios developed by Kurylyk and
403 MacQuarrie (2013) to simulate groundwater discharge to streams in the unconfined
404 granular aquifer of Otter Brook (NB) under climate change. Groundwater flow, freezing,
405 and thawing were simulated with the SUTRA model, for two aquifer configurations
406 (aquifer discharge to the brook or to lateral seeps). The summer groundwater discharge
407 rates varied between -6% and +39%, with an increase of discharging water temperature

408 up to 3.6°C. The authors conclude that small and shallow aquifers are susceptible to air
409 temperature increases and that thermal refugia could be impacted through warmer
410 groundwater inflows.

411 In the Richibucto region of New Brunswick, Green and MacQuarrie (2014) used the
412 SEAWAT and HELP models to examine relative impacts of climate change-induced
413 variations in recharge, sea level rise and increased groundwater extraction on saltwater
414 intrusion (2011-2100). Two recharge scenarios were based on a previous study (using the
415 HELP model; Jacobs 2011): one scenario reflected the projected changes developed by
416 Jacobs (2011), and the second doubled the percent change of those projections relative to
417 historic conditions. Recharge changed between -27% and -5%. The impact of decreasing
418 recharge was the most important at depths less than 60 m below sea level. Interestingly,
419 sea-level rise had the least important effect on seawater intrusion in shallow to
420 intermediate aquifers for the future scenarios. Because of the importance of both recharge
421 and pumping on seawater intrusion, the authors suggested that actions to control land use
422 influencing recharge, as well as pumping rates, may help to protect coastal fresh-
423 groundwater supplies.

424 Rivard et al. (2014) investigated the impacts of climate change on the HELP-simulated
425 recharge for the Annapolis Valley (NS) (546 km²). The results showed a change in
426 annual recharge from +14 to +45%. Recharge changes during the growing season (May
427 to October) varied between -33% and -4%.

428 Lemieux et al. (2015) simulated the depth and shape of the groundwater transition
429 zone between freshwater and seawater in the Magdalen Islands (Québec; 200 km²). The
430 simulation of density-dependent flow was performed with the FEFLOW model along a

431 vertical 2D cross section. The authors calculated recharge values using a surface water
432 budget and the climate scenario variables, and in the model imposed the recharge to
433 evolve linearly from current conditions to the worst-case recharge scenario (-30%). In
434 these conditions, the impact of sea level rise was larger than that of both coastal erosion
435 and reduced recharge on the position of saltwater–freshwater interface which could
436 migrate inland over a distance of 37 m.

437 Paradis et al. (2016) investigated how nitrate concentrations in groundwater might
438 evolve under climate change conditions and with changes in agricultural practices in
439 Prince Edward Island (PEI) (5,660 km²). The HELP model was used to simulate recharge
440 with climate scenarios used with a groundwater flow model developed in FEFLOW.
441 Nitrate concentrations resulting from residual soil nitrate for eight scenarios of
442 agricultural changes were simulated with the four recharge scenarios. The results showed
443 changes in annual recharge that ranged between -12% and +7%. The generally increasing
444 nitrate concentrations were primarily attributed to the attainment of steady-state
445 conditions under present-day nitrogen loading, and to an increase in nitrogen loading in
446 some agricultural scenarios. Only 0% to 6% of the increase in nitrate concentrations in
447 groundwater was explained by changes in the recharge scenarios.

448

449 **DISCUSSION**

450 *Changes in recharge*

451 Comparing recharge rates between the different studies is not an easy task since recharge
452 is sometimes reported as mean annual values and sometimes as seasonal values. Some
453 studies simulate infiltration towards deeper soil (aquifer) layers, but do not explicitly

454 quantify recharge (Quilbé et al. 2008; Levison et al. 2014b; CEHQ 2015). The studies
455 based on groundwater flow models (Bourgault et al. 2014; Green and MacQuarrie 2014;
456 Levison et al. 2014; Lemieux et al. 2015) have simulated worst-case recharge scenarios
457 and present the largest recharge decreases (maximum annual recharge decrease of 50%,
458 Bourgault et al. 2014). Although these values are useful to plan for extreme conditions,
459 they do not necessarily reflect the complete array of possible future conditions.

460 Five out of six studies based on the HELP model show recharge increases (Brouwers
461 2007; Jyrkama and Sykes 2007; Kurylyk and MacQuarrie 2013; Rivard et al. 2014;
462 Motiee et al. 2017) while four out of six studies with integrated models show at least one
463 scenario with increased recharge (Colautti 2010; Sulis et al 2011; Sulis et al 2012;
464 Cochand 2014), and four out of six studies with integrated models show at least one
465 scenario with decreased recharge (Colautti 2010; Sultana and Coulibaly 2011; Sulis et al.
466 2012; Cochand 2014). Even though it is generally calibrated on total flow and baseflow
467 separation, HELP simulates water that percolates below the root zone. It is often not clear
468 whether or not this water reaches the aquifer and travel a significant distance with the
469 saturated zone. This could explain why, when these processes are considered in
470 integrated models where lateral and vertical aquifer heterogeneity is included, the
471 changes in recharge are less conclusive.

472 The reviewed studies do not show any clear trend from West to East. Apart from the
473 type of model used in the simulations, the discrepancies between studies could be due to
474 the climate models, to the post-processing methods, or to the projected horizons.
475 However, no clear causes could be identified with the available studies. The variability
476 introduced by these components of a study have been identified by Quilbé et al. (2008)

477 and Kurylyk and MacQuarrie (2013). For example, GCMs poorly simulate precipitation
478 but statistical downscaling can improve the projected values and might be superior to the
479 delta change method since it can include changes in rainfall occurrence (Quilbé et al.
480 2008) which could impact recharge fluxes.

481 The studies generally agree that a global warming would reduce snow accumulation
482 during the winter. More frequent episodes of warmer temperatures, less snow
483 accumulation, and rain during winter are expected to increase winter recharge and lower
484 spring recharge. In the future projections, recharge is often higher in the winter and lower
485 in the spring season (Brouwers 2007; Colautti 2010; Sulis et al. 2011; Sultana and
486 Coulibaly 2011; Cochand 2014; CEHQ 2015). Sulis et al. (2012) also determined that
487 changes in the duration of the wet season had a large impact on recharge. However, the
488 dynamics of winter recharge still need to be investigated to fully understand how it will
489 be impacted by climate change.

490 The three studies that focused on coastal groundwater resources in similar geological
491 formations have shown contrasting results. Those of Hansen (2012) and Green and
492 MacQuarrie (2014) indicate that sea level rise would have a limited impact on salt water
493 intrusion, while that of Lemieux et al. (2015) showed that sea level rise would have the
494 largest impact (compared to erosion and change in recharge). Lemieux et al (2015) report
495 that when using a time frame similar to that of Green and MacQuarrie (2014), the impact
496 of sea level rise becomes more important in the Magdalen Island study. It should be
497 acknowledged that sea level rise can be locally variable, so its impacts on seawater
498 intrusion are expected to exhibit more spatial patterns than recharge changes. This
499 underlines an important component of climate change impact studies, i.e. that aquifers

500 can react slowly to changes in recharge and other boundary conditions. Long-term and
501 continuous transient-state simulations are expected to better reflect the slow response of
502 aquifer reservoirs.

503 The studies also show that more inter-annual variability in recharge should be
504 expected during the summer, due to warmer air temperature and to more intense
505 precipitation generating runoff instead of recharge (not easily captured by models that are
506 based on a daily time step; Allen et al. 2014). This could induce longer drought periods
507 affecting the summer baseflows (Brouwers, 2007; Quilbé et al. 2008; Sulis et al. 2011;
508 Sultana and Coulibaly 2011; Cochand 2014; Levison et al. 2014a; CEHQ 2015). Also, if
509 recharge occurs earlier in the spring, small aquifers and those located in headwater basins
510 may be fully drained by summer months which would induce early low baseflow
511 conditions. Since municipal and agricultural users generally need more groundwater
512 during the summer period than during the winter, these changes could greatly affect
513 economic activities. This underlines the importance of representing soil and hillslope
514 storage variations through time and the importance of changing sequences of rainy days
515 on the simulation of recharge (Sulis et al. 2012). Again, this stresses the importance of
516 using long-term transient-state simulations.

517

518 ***Changes in groundwater discharge to surface water***

519 Only one Ontario-based study (out of the three that reported baseflow values) showed
520 dominating conditions of decreased baseflows (McLaren and Sudicky 1993) while that of
521 Colautti (2010) showed mixed conditions. In Québec, generally decreasing baseflows
522 were reported (Quilbé et al. 2008; Sulis et al. 2011; Sulis et al 2012; Bourgault et al.

523 2014; Cochand 2014; CEHQ 2015), but mixed conditions were also reported (Levison et
524 al. 2014a; 2014b). The only study in Atlantic Canada that reported baseflows (Kurylyk et
525 al. 2014) shows mainly increasing baseflows for a small granular aquifer.

526 Sulis et al. (2011) and Levison et al. (2014a) clearly identified that changes in local
527 aquifers near surface water bodies could lead to more frequent reversals of the hydraulic
528 gradient between aquifers and surface water bodies. These reversals lead to reductions in
529 baseflow and adds pressure on ecological habitats and water users in streams and ponds
530 during droughts and low water periods in summer (e.g., Levison et al. 2014b). The
531 absence of significant decreases in baseflows under climate change conditions in New
532 Brunswick can appear counterintuitive since Rivard et al. (2009) have identified
533 decreasing trends in baseflows in past time series for Atlantic Canada. However, the only
534 climate change impact study reported here (Kurylyk et al. 2014) concerns a very small
535 watershed which might not be representative of larger scale conditions.

536 Although early studies of the impact of climate change on water resources did not
537 include consideration of groundwater flow (e.g., Southam et al. 1999), the most recent
538 studies using surface flow models incorporate at least a simplified representation of
539 aquifer reservoirs (e.g., Quilbé et al. 2008; CEHQ 2015). In these cases, the surface flow
540 model simulates baseflows which can be attributed to a groundwater contribution to the
541 river and changes in these flows under climate change conditions can be studied.
542 Interestingly, four out of five applications of integrated models show at least one scenario
543 with baseflow decrease while two out of five applications show at least one scenario of
544 baseflow increase. It is self-evident that integrated surface water-groundwater models
545 simulate more completely and probably more reliably baseflow conditions. However, in

546 this review, the baseflow results appeared to be independent of the type of model
547 (groundwater flow or integrated surface-groundwater flow model). They also appeared to
548 be independent of the climate model, and of the post-processing method.

549

550 *Climate change impacts in conjunction with other pressures*

551 Agricultural, urban and potable water pressures on groundwater availability are
552 already of concern in most of the inhabited regions in Eastern Canada. Climate change is
553 expected to influence indirectly groundwater use which can evolve through climate-
554 driven or socio-economically-driven land use change (Taylor et al. 2012). Water quantity
555 stress assignments have been performed on the Grand River watershed and elsewhere to
556 investigate the effect of possible future increases of water use on water availability (e.g.,
557 AquaResources 2009a; 2009b) and water use conflicts (e.g., Lavigne et al. 2010a;
558 2010b). Kurylyk and MacQuarrie (2013) proposed that when studying recharge on a
559 projected horizon longer than a few decades, land use changes and socio-economic
560 factors should be taken in consideration. Nikolic and Simonovic (2015) (not reported in
561 detail herein because the paper did not provide quantified changes in recharge or
562 baseflows) have shown that at the sub-watershed scale, increasing permits to take water
563 for agriculture could have an adverse impact on the groundwater resources. Bourgault et
564 al. (2014) provide similar results and showed that increased groundwater pumping could
565 have a larger impact on groundwater resources than decreased recharge in a St. Lawrence
566 Lowlands granular superficial aquifer. Hansen (2012) and Green and MacQuarrie (2014)
567 have shown that groundwater use through pumping in coastal aquifers can exacerbate
568 saltwater intrusion. Analytical solutions have been developed to better apprehend these
569 conditions and help water managers (e.g., Ferguson and Gleeson 2012).

570 Changing land use through deforestation, expansion or changes in agricultural
571 activities or urban areas, and drainage of wetlands can have impacts on surface and
572 subsurface hydrology (Mishra et al. 2010). For example, Fossey et al. (2016) have used
573 surface flow modeling to show that isolated wetlands located in the upper part of a
574 watershed have a larger effect on maintaining low flows and damping high flow than
575 wetlands located downstream. If wetland drainage is combined with a dryer or flashier
576 future climate, the impacts of groundwater and surface water resources could be
577 exacerbated. Paradis et al. (2016) have also shown that changes in nitrogen loading that
578 could be a consequence of increased temperatures leading to different crop choices, or
579 that could result from the intensification of agricultural activities, would have a larger
580 impact on groundwater nitrate concentrations than changes in recharge.

581

582 *Modelling at different scales*

583 The reviewed studies cover a large range of scales. Local scale studies ($< 100 \text{ km}^2$)
584 (Brouwer 2007; Hansen 2012; Kurylyk and MacQuarrie 2013), usually need strong
585 collaborative support from specialists familiar with the regional geology and the local
586 characteristics (Frey et al. 2016). Data is often most available at the watershed scale (100
587 to $1,000 \text{ km}^2$) where watershed organizations contribute to knowledge acquisition and
588 data availability. At that scale, most of the studies of the last decade have been
589 implemented with integrated models (Colautti 2010; Sultana and Coulibaly 2011; Sulis et
590 al. 2011; 2012; Cochand 2014; Levison et al. 2014b). Groundwater flow models are still
591 used (Levison et al. 2014a; Kurylyk et al. 2014; Lemieux et al. 2015; Levison et al.
592 2016), but they necessitate simplifications in the recharge processes or the coupling of a
593 recharge model to a groundwater flow model (e.g., Kurylyk et al. 2014; Paradis et al.

594 2016). This has unquantified impacts on the simulated conditions and the surface-
595 unsaturated zone-saturated aquifer feedback processes.

596 At the regional scale (2,000 to 10,000 km²) hydrogeological data and detailed
597 descriptions of aquifer properties do not always exist. Major funding has been invested in
598 aquifer characterization in the Province of Québec in the last decade (see MDDELCC
599 2017a for a full list of reports since 2009). In Ontario, there have been many regional
600 scale hydrogeological studies since the 1980s and 1990s (e.g., Howard and Beck 1986;
601 Novakowski and Lapcevic 1988; Sharpe et al. 1996; Rudolph et al. 1998). These detailed
602 groundwater data over a large region, such as in the Grand River watershed in Ontario,
603 allowed researchers to develop complete groundwater description and integrated models
604 (Jyrkama and Sykes 2007; Colautti 2010). In Atlantic Canada, the Canadian Geological
605 Survey has performed regional hydrological characterization studies in the last decade as
606 part of the Canadian Groundwater Inventory (Paradis et al. 2007; Rivard et al. 2008;
607 Rivard et al. 2012). All these initiatives have led to the development of databases that can
608 support model development for climate change studies.

609 Climate change impacts on the groundwater dynamics in Eastern Canada have not yet
610 been investigated on a provincial scale. Natural Resources Canada has recently initiated
611 the development of a fully-integrated groundwater-surface water, climate impact model
612 for the southern Ontario Phanerozoic Basin Region (Frey et al. 2016; NRCan 2017). The
613 Québec Ministry of Environment (MDDELCC) has recently initiated the development of
614 an integrated groundwater-surface water model for southern Québec that will be used to
615 better understand the impact of climate change on water resources. These models will
616 facilitate anticipating long-term changes over large areas, including those resulting from

617 climate change. This is clearly a positive development towards integrated water
618 management and adaptation for future conditions. These models could include other
619 cumulative stresses occurring on groundwater resources, such as changes in land use and
620 increasing pumping for drinking water, industrial or irrigation purposes. The combination
621 of studies and models existing at different scales, developed with various purposes in
622 mind, will truly aid decision making for groundwater management.

623

624 *Uncertainty in future conditions*

625 The large range of possible future recharge conditions in Eastern Canada is not
626 uncommon in climate change studies. It has been reported elsewhere and appears to be
627 intrinsic to the study of climate change effects (e.g., Green et al. 2011). This variability
628 can be due to the use of various climate models, emissions scenarios, data treatment
629 methods (e.g., downscaling), and future time horizons. Although some studies have
630 aimed specifically at better understanding this component and argue that using different
631 downscaling methods and different sources of data is a necessity (Quilbé et al. 2008), the
632 studies reviewed here did not allow to identify which of these methods should be
633 prioritized over others.

634 Variability in the results also comes from the flow models themselves, notably from
635 over-simplification of the geological conditions stemming from insufficient
636 hydrogeological data used to build the model. The type of modeling approach to be used
637 is also crucial to the array of possible future conditions. The review presented here
638 indicates that coherent results appear between the three regions when similar modeling
639 approaches are used. Among these, and rather intuitively, the integrated surface-

640 groundwater flow models may be more robust to simulate climate change impacts. They
641 allow a more holistic understanding of the entire water cycle and of the feedbacks
642 between reservoirs, within a single mathematical framework, thus alleviating time and
643 spatial scale errors. Going one step further, Sulis et al. (2017) report an application of
644 coupling water flow, vegetation and atmospheric processes in western Germany. This
645 type of integrated approach is expected to become increasingly used as computer
646 capacities continue to increase in the next decade.

647 Using short time series of heads and flow rate data for model calibration data can also
648 be responsible for model uncertainty because they do not include a wide array of possible
649 meteorological and hydrological conditions (Moeck et al. 2016). A model that has been
650 calibrated based on years of wet conditions is not necessarily robust to simulate dry
651 conditions, or a succession of wet and dry periods. This is certainly a challenge, and
652 although most Canadian provinces have a reasonably well maintained surface water
653 monitoring network, groundwater monitoring has only recently received serious
654 attention. Recent efforts have been invested in Québec since 2000 to install a province-
655 wide groundwater monitoring network (MDDELCC 2017b). Ontario has had the
656 Provincial Groundwater Monitoring Network since 2001, which has 492 monitoring
657 points across the province (MOECC 2017). In Nova Scotia, the Groundwater
658 Observation Well Network was established in 1965 and currently hosts 40 observation
659 wells (Government of Nova Scotia, 2017). In New Brunswick, the groundwater
660 observation well network was established in the early 1970s, was disbanded in 2000 and
661 is currently being re-instrumented with a limited number of stations. In Prince Edward
662 Island, groundwater elevations are available for 14 monitoring wells, with data from as

663 early as the 1967 for some wells available online (Government of Prince Edward Island,
664 2017).

665

666 ***Recommendations***

667 The first recommendation from this review is that multiple scenarios from an array of
668 climate models and different climate change scenarios should be used. This follows
669 recommendations by Holman et al. (2011), and was observed in most of the reviewed
670 studies. The most recent studies use RCP scenarios and span future horizons that reach
671 2100. It is now relatively easy for researchers, consultants and water managers who wish
672 to perform climate change impact studies to have access to databases that provide
673 downscaled climate data across Canada. For example, the Pacific Climate Impacts
674 Consortium provides temperature and precipitation data from statistically downscaled
675 climate scenarios (on a grid of approximately 10 km resolution) for the entire country for
676 1950-2100 ([https://www.pacificclimate.org/data/statistically-downscaled-climate-](https://www.pacificclimate.org/data/statistically-downscaled-climate-scenarios)
677 [scenarios](https://www.pacificclimate.org/data/statistically-downscaled-climate-scenarios)). The Climate Change Data Portal (<http://ccdp.network/>) also provides
678 dynamically downscaled climate change scenarios (temperature and precipitation) for
679 Canada and other regions of the world. The Ouranos Consortium provides on demand
680 climate scenarios based on regional or global climate models and is currently developing
681 a web-access platform to make climate scenarios publicly available (*PAVICS – Power*
682 *Analytics and Visualization for Climate Science*).

683 The decision of which model to use is often not an easy one since it depends on
684 available data, available time to perform the study and local expertise. Nevertheless,
685 some insights arise from this review. It is clear that surface flow models provide only

686 limited insight to the impact of climate change on groundwater resources because they
687 use highly simplified representations of aquifers and generally not calibrated to fully
688 represent baseflow conditions. They are thus not the preferred tool when studying climate
689 change impacts on groundwater resources. Groundwater flow models are useful to
690 understand conditions in specific areas, but they are tributary to the separate simulation of
691 recharge processes and this adds some level of uncertainty in the simulation of climate
692 change impacts. Integrated surface water-groundwater flow models are a logical
693 approach to understand the impacts of climate change on water resources, and the
694 feedbacks between different water reservoirs. The second recommendation is that these
695 models should be favored when possible in future studies.

696 This review also highlighted the important spatial variability in simulated changes for
697 recharge and baseflows. This variability can only be taken into account by using models
698 at different scales, depending on the issues under consideration. Local models, watershed
699 models, regional or supra-regional models all contribute to better understanding this
700 variability and to provide managers with decision-making tools. Topographic location,
701 geological conditions have been identified as having a possible influence on how aquifers
702 will respond to climate change. This review also has underlined the importance of
703 performing long term transient-state simulations to assess the impact of temporal
704 variability, as well as the long-term storage potential of aquifers. The third
705 recommendation is that model efforts should represent long-term climate change impacts
706 at different scales.

707 Other anthropogenic pressures such as land use changes and increased pumping rates
708 are also of crucial importance to assess combined effects and cumulated impacts. The

709 fourth recommendation is that these should be included in future studies to provide a
710 background of information to assess tipping points for groundwater resources and
711 ecosystems. The case of cumulated impacts on coastal groundwater resources has been
712 made in this review. The impact of climate change on groundwater dependent ecosystems
713 subjected to land development pressure and contamination should also be further
714 investigated.

715 Spatial coverage of the territory of Eastern Canada is another important issue. The
716 studies of climate change impacts on groundwater resources have yet focused only in the
717 southern part of Eastern Canada. The fifth recommendation is that more studies should be
718 performed to assess the impact of climate change in Northern Ontario, North of the St.
719 Lawrence River, and in the Bas-Saint-Laurent and Gaspésie regions, as well as in
720 Newfoundland and Labrador. Additional studies would also be useful to better
721 understand how salt water intrusion will be affected by sea level rise, changes in recharge
722 and pumping in coastal regions of Québec and Atlantic Canada. Locations/regions for
723 future studies should be prioritized based on criteria such as expected increases in water
724 supply demand, ecological flow requirements, resource extraction, and other issues
725 relevant to provincial and municipal governments. This may shift as governmental
726 priorities evolve in each province. Stakeholders who also often provide funding for this
727 type of research (e.g., provincial environment and agricultural ministries, in consultation
728 with other interested parties) can continue to set priorities for critical locations for future
729 investigation based on the above criteria.

730

731 **CONCLUSION**

732 This paper aimed to review the state of knowledge about simulated climate change
733 impacts on groundwater dynamics in Eastern Canada, including the provinces of Ontario,
734 Québec, New Brunswick, Nova Scotia, and Prince Edward Island. Three major issues
735 have been highlighted from these studies: i) no overall trend in time or space in recharge
736 could be identified from the reviewed studies, but for Eastern Canada, but more inter-
737 annual variability throughout the year is expected due to seasonal shifts in recharge; iii)
738 groundwater discharge to surface water bodies tends to be reduced in future scenarios,
739 and particularly during summer droughts. This exercise has provided a valuable reflection
740 of our current understanding of possible future conditions. The analysis contributed to
741 identifying how to improve climate change impact studies in such a way that they can be
742 more useful for water managers.

743 There is clearly a need to establish guidelines for performing climate change impact
744 studies on groundwater resources. This review provided new insights that lead to the
745 following five recommendations for future studies: 1) use a variety of climate models and
746 emission scenarios; 2) promote the use of integrated models when possible; 3) study
747 long-term climate change impacts on groundwater resources at different scales; 4)
748 simulate the combined effects of climate change and other pressures; and 5) develop
749 models that cover other regions of Eastern Canada as dictated by stakeholders and water
750 managers. It is clear that a good understanding of aquifer geometry and groundwater
751 flow dynamics, a dense coverage of long-term monitoring stations for piezometric heads
752 and river flow rates, and the development of integrated models that are maintained in the
753 long-term would facilitate water management and planning in a changing climate.

754

755

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1052 **Figure caption**

1053 Figure 1. a) Location of all the available studies reporting climate change impacts on
1054 groundwater resources in Eastern Canada (the dots correspond to the center of the
1055 study area), b) Ontario-based studies, c) Québec-based studies, and d) Atlantic Canada
1056 studies

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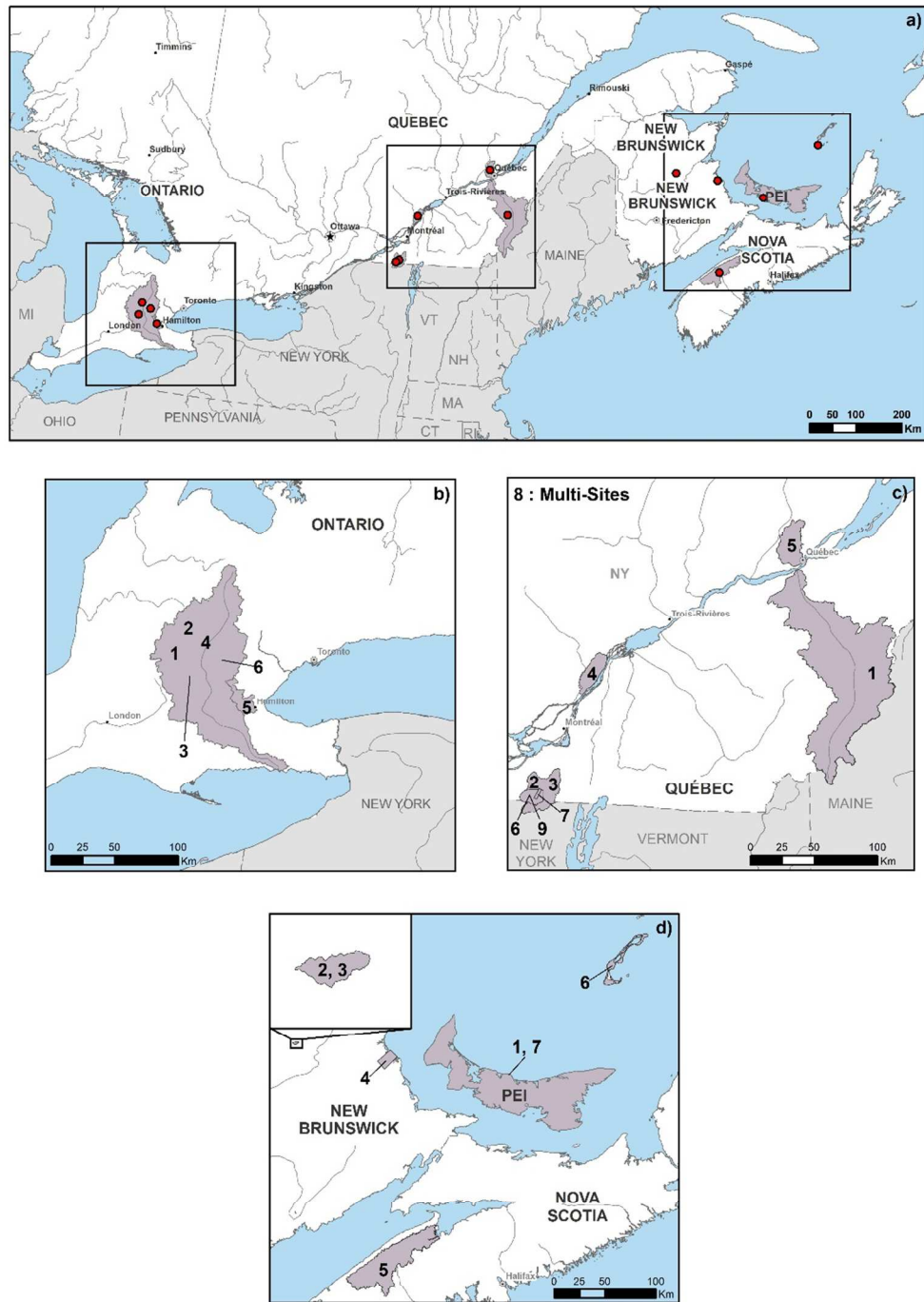


Figure 1
Larocque et al. 2017
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Table 1. Studies on climate change impacts on groundwater for the Province of Ontario

No.	Study area	Size (km ²)	Model	Post-processing method(s)	Emission scenario(s)	Projected horizon(s)	Simulated change in baseflow	Simulated change in recharge
1.	Subregion of Grand River (ON)	Not stated	2D steady-state groundwater flow model	Delta change method	2 X CO ₂	2050	-39% to -17%	-15% to +35%
2.	Grand River (ON)	6,800	HELP	Inverse distance squared	General predictions from IPCC (2000)	40 years	n.a.	+10% to +53%
3.	Alder Creek - Grand River (ON)	80	HELP HydroSphere	Delta change method (monthly) ¹	A2x, B2x	2020-2080	n.a.	+0.36 mm to +4.12 mm
4.	Grand River (ON)	6,800	HydroGeoSphere	Perturbed historical records	n.a. ²	40 years	-15% to +59%	-5% to 22%
5.	Spencer Creek (ON)	291	MikeSHE	Statistical downscaling Time-lagged-forward neural networks	A2	2046–2065	n.a.	-6% to -0.5%
6.	Guelph region of the Grand River (ON)	Not stated	HELP	Delta change method (monthly)	n.a.	2010-2050	n.a.	+7 to +12%

1. McLaren and Sudicky (1993)

3. Brouwers (2007)

5. Sultana and Coulibaly (2011)

2. Jyrkama and Sykes (2007)

4. Colautti (2010)

6. Motie and McBean (2017)

¹Scaling factors were available (monthly) for 2020, 2050 and 2080 and linear interpolation was used between the time slices to distribute monthly scaling factors to a daily timestep.

²Synthetic scenarios were constructed based on modifying the 1960-1999 precipitation record (-5%, +5%, +10%, +15%, +20%), bounded by GCM-based climate scenarios.

³The values are average variations in mm per month (the minimum is for urban land use and the maximum is for agricultural areas) since the reference values were not provided per land use type by the author for the entire study area.

⁴A weather generator algorithm applied changes to a local-climate time series to create future climate conditions.

Table 2. Studies on climate change impacts on groundwater for the Province of Québec

No.	Study area	Size (km ²)	Model	Post-processing method(s)	Emission scenario(s)	Projected horizon(s)	Simulated change in baseflow	Simulated change in recharge
1	Chaudière River (QC)	6,682	HYDROTEL	Delta change method (monthly) based on GCM variables; statistical downscaling; combination of the two methods.	A2, B1, B2	2010-2039	-29% to -5% ¹	n.a.
2	des Anglais River (QC)	690	CATHY	Delta change method (monthly) based on CRCM data	A2	2038-2070	-56% ²	+16% ³
3	des Anglais River (QC)	690	CATHY	Dynamic downscaling (for the RCM models) and regridding on a 50 km resolution grid (weighed inverse distance)	A2	2041-2065	Increased low flow occurrence	-15% to +4% ⁵
4	Lanoraie peatland (QC)	364	MikeSHE (only groundwater flow)	Delta change method	A1B, A2, B1	2040-2069	-41% to -16%	-50% to -20% ⁶
5	Saint-Charles River (QC)	553	HydroGeoSphere	Delta change method (monthly)	A1B, A2, B1	2071-2100	-25% to -10%	-50% (spring) +150% to +300% (winter) -8% to -4% (summer)
6	Covey Hill Natural Lab. (QC)	173	MODFLOW	Dynamic downscaling	A1B, A2	2041-2070	-44% to +14%	-30% to +10% ⁵
7	Covey Hill Natural Lab. (QC)	173	HydroGeoSphere	Dynamic downscaling	A1B, A2	2041-2070	+5% to +6% ⁷	n.a.
8	Multi-sites	500 to 20,000	HYDROTEL	Delta quantile mapping	RCP4.5, RCP8.5	2041-2070	-63% to -6%	n.a.
9	Covey Hill Natural Lab. (QC)	173	MODFLOW	Dynamic downscaling	A1B, A2	2041-2070	+10% to +14% ⁸ 18% ⁹	-4% to +15% ⁶

1. Quilbé et al. (2008)

3. Sulis et al. (2012)

5. Cochand (2014)

7. Levison et al. (2014b)

9. Levison et al. (2016)

2. Sulis et al. (2011)

4. Bourgault et al. (2014)

6. Levison et al. (2014a)

8. CEHQ (2015)

¹: Only the scenarios based on the Delta method showed decreases in baseflows, the scenarios based on statistical downscaling showed no changes in baseflows.

²: Reduction in summer (June, July, and August) flows at the outlet (Figure 7 in Sulis et al. 2011).

³: Annual variation in total recharge over the watershed (Figure 8 in Sulis et al. 2011).

⁵: 11 out of the 12 simulated futures projected a decreasing annual recharge.

⁶: Recharge scenarios were imposed on the groundwater flow model based on a surface water budget calculated with the climate scenarios.

⁷: Flow rate increase in the springs depending on the altitude, considered here as baseflow.

⁸: Range in average baseflow changes for three rivers.

⁹: Average flow rate increase for the simulated springs, considered here as baseflow.

Table 3. Studies on climate change impacts on groundwater for the Atlantic Provinces

No	Study area	Size (km ²)	Model	Post-processing method(s)	Emission scenario(s)	Projected horizon(s)	Simulated change in baseflow	Simulated change in recharge
1	Summerside (PEI)	4.9	SEAWAT (MODFLOW)	Dynamic downscaling	A1B, A2, B2, multi-ensemble approach	2100	n.a.	+5% to +8% ¹
2	Otter Brook (NB)	9.5	HELP	Delta change method (daily), Hybrid multiple regression, and Dynamical downscaling	A1B, A2, B1	2046-2065	n.a.	-6% to +58%
3	Otter Brook (NB)	9.5	SUTRA	Delta change method (daily), Hybrid multiple regression, and Dynamical downscaling	A1B, A2, B1	2046-2065	-6% to +31% ²	-6% to +58% ³
4	Richibucto (NB)	142	HELP SEAWAT (MODFLOW)	Mean deviations from climate indices for 2020s, 2050s and 2080s calculated	A1B, A2	2011-2100	n.a.	-27% to -5%
5	Annapolis Valley (NS)	546	HELP	Dynamic downscaling, and Monthly delta method	A2	2041-2070	n.a.	+14% to +45% ⁴
6	Magdalen Islands (QC)	200	FEFLOW	Delta change method (monthly)	A1B, A2, B1	2013-2040	Saltwater-freshwater interface migrates inland by 37 m.	-30% to 0% ⁵
7	Prince Edward Island (PEI)	5,660	HELP FEFLOW	Statistical downscaling	A2, B2	2040-2069	n.a.	-12% to +7%

1. Hansen (2012)

3. Kurylyk et al. (2014)

5. Rivard et al. (2014)

7. Paradis et al. (2016)

2. Kurylyk and MacQuarrie (2013)

4. Green and MacQuarrie (2014)

6. Lemieux et al. (2015)

¹: Assuming that recharge will change with the same percentage as precipitation from the climate change scenarios.²: Values for summer low flows.³: Imposed from Kurylyk and MacQuarrie (2013).⁴: Mean annual values from the seven simulated scenarios.⁵: Recharge was made to evolve linearly from current conditions (0% change) to the worst-case scenario recharge obtained from water budget calculations.

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